# Surprising, but not unexpected, multi-decadal pause in Arctic sea ice loss

Mark England<sup>1</sup>, Lorenzo M Polvani<sup>2</sup>, James A Screen<sup>1</sup>, and Anthony ChunYin Chan<sup>1</sup>

<sup>1</sup>University of Exeter <sup>2</sup>Columbia University

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#### Abstract

Over the past two decades, Arctic sea ice loss has slowed considerably, with no statistically significant decline in September sea ice area since 2005. This pause is robust across observational datasets, metrics, and seasons. Large-ensemble CMIP5 and CMIP6 simulations reveal that such periods with no sea ice decline under increasing greenhouse gas emissions are not unusual. Analysis of ensemble members that simulate analogues of the observed pause indicates that the current slowdown could plausibly persist another five to ten years. The modelling evidence suggests that internal variability has substantially offset anthropogenically forced sea ice loss in recent decades, although possible contributions from changes in the forced response remain uncertain. Overall, this observed pause in Arctic sea ice decline is consistent with simulated internal variability superimposed on the long term trend according to the bulk of the climate modelling evidence.

#### Surprising, but not unexpected, multi-decadal pause in 1 Arctic sea ice loss 2

M. R. England<sup>1</sup>, L. M. Polvani<sup>2,3</sup>, J. Screen<sup>1</sup>, and A. C. Chan<sup>1</sup>

<sup>1</sup>Department of Mathematics and Statistics, University of Exeter, UK <sup>2</sup>Lamont-Doherty Earth Observatory, Columbia University, New York, USA <sup>3</sup>Department of Applied Physics and Applied Mathematics, Columbia University, New York, USA

## Key Points:

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8	•	The loss of Arctic sea ice cover has undergone a pronounced slowdown over the
9		past two decades, across all months of the year.
10	•	Rather than being an unexpected rare event, comprehensive climate models from
11		CMIP5 and CMIP6 simulate such pauses relatively frequently.
12	•	According to these climate model simulations, this pause in the loss of Arctic sea
13		ice could plausibly continue for the next 5-10 years.

Corresponding author: Mark Ross England, m.england2@exeter.ac.uk

#### 14 Abstract

Over the past two decades, Arctic sea ice loss has slowed considerably, with no statis-15 tically significant decline in September sea ice area since 2005. This pause is robust across 16 observational datasets, metrics, and seasons. Large-ensemble CMIP5 and CMIP6 sim-17 ulations reveal that such periods with no sea ice decline under increasing greenhouse gas 18 emissions are not unusual. Analysis of ensemble members that simulate analogues of the 19 observed pause indicates that the current slowdown could plausibly persist another five 20 to ten years. The modelling evidence suggests that internal variability has substantially 21 offset anthropogenically forced sea ice loss in recent decades, although possible contri-22 butions from changes in the forced response remain uncertain. Overall, this observed pause 23 in Arctic sea ice decline is consistent with simulated internal variability superimposed 24 on the long term trend according to the bulk of the climate modelling evidence. 25

#### <sup>26</sup> Plain Language Summary

Over the last 20 years, the decline of Arctic sea ice has slowed down substantially. 27 Climate models (from CMIP5 and CMIP6) show that pauses in sea ice loss across mul-28 tiple decades can happen, even as greenhouse gas emissions continue to rise. When we 29 compare the current slowdown to similar pauses in model simulations, we see that this 30 "hiatus" could plausibly continue for another five to ten years. Most of the evidence from 31 these climate models suggests that natural climate variations have played a large part 32 in slowing the human-driven loss of sea ice. However, it is not entirely certain whether 33 changes in the human influence on climate (the "forced response") have also contributed. 34 Overall, while it may sound surprising that Arctic sea ice loss has slowed down even as 35 global temperatures hit record highs, the climate modelling evidence suggests we should 36 expect periods like this to occur somewhat frequently. 37

#### 38 1 Introduction

The loss of Arctic sea ice over the past half century is one of the most clear and 39 well-known indicators of human-induced climate change (IPCC, 2021; Copernicus, 2024). 40 September sea ice area has nearly halved since the beginning of the satellite era in 1979 41 (Fetterer et al., 2017; Stroeve & Notz, 2018), and during the same period, estimated Arc-42 tic sea ice volume has decreased by over 10,000 km<sup>3</sup> (Kwok, 2018). Record-breaking sum-43 mer sea ice minimums in 2007 (Stroeve et al., 2011) and 2012 (Parkinson & Comiso, 2013; 44 J. Zhang et al., 2013) fuelled predictions, which with hindsight look overly alarmist, that 45 the Arctic would experience its first ice-free summer before 2020 (Maslowski et al., 2012; 46 Wadhams, 2016). Adding to this, the Arctic has been warming up to four times faster 47 than the global average (Rantanen et al., 2022). It has been further proposed that global 48 warming might be accelerating, culminating in record breaking warmth in recent years 49 (Samset et al., 2023; Hansen et al., 2025; Merchant et al., 2025). As Arctic sea ice cover 50 is strongly tied to global temperatures (Notz & Stroeve, 2016), there would be little ex-51 pectation of a multi-decadal *slowdown* in Arctic sea ice loss. And yet, as we will show, 52 such a slowdown has been occurring in the last two decades. 53

Recall that, over the past century, periods of increasing anthropogenic greenhouse 54 emissions without sustained sea ice loss - the mid-20th century (Walsh et al., 2017) - have 55 already occurred. From the 1940s to the 1970s Arctic sea ice cover expanded (Gagne et 56 al., 2017), with the largest increases in the Chukchi, East Siberian, Laptev, Kara and 57 Barents Seas. However, anthropogenic forcing in the mid-20th century was very differ-58 ent compared to the one of the past two decades. Industrial aerosol emissions from Eu-59 rope and North America contributed substantially to the positive multi-decadal trend 60 in Arctic sea ice area and associated Arctic cooling in the mid-20th century (Fyfe et al., 61 2013; Nafaji et al., 2015; Gagne et al., 2017; England et al., 2021); but these aerosol sources 62 are far smaller today (Szopa et al., 2013; Lund et al., 2019). However, when anthropogenic 63

aerosols are being discussed in the context of the recent past, it is with regards to the
phase out of aerosol emissions from ship tracks which have potentially contributed to enhanced global warming since 2020 (Manshausen et al., 2022; Yoshioka et al., 2024). In
fact, Yoshioka et al. (2024) find that the simulated warming response to these reduced
sulphur emissions is largest in the Arctic. So, the lessons of the past may not be a reliable guide for understanding current trends.

It is important to appreciate that the observed trend in Arctic sea ice cover over 70 a given period is composed of a contribution caused by anthropogenic emissions, denoted 71 72 the forced response, and a contribution from unforced fluctuations associated with internal climate variability (England et al., 2019; England, 2021; Dörr et al., 2023; Shen 73 et al., 2024). Anthropogenically-forced changes which may contribute to a reduction in 74 Arctic sea ice loss over the past two decades include a forced slowdown in the Atlantic 75 Meridional Overturning Circulation (Lee & Liu, 2023), and changes in the emissions from 76 biomass burning, both in the magnitude (Blanchard-Wrigglesworth et al., 2025), and the 77 variability (DeRepentigny et al., 2022). One would imagine, however, that the aforemen-78 tioned reduction of sulphur emission from shiptracks (Yoshioka et al., 2024) would lead 79 to an acceleration rather than a deceleration of sea ice loss since 2020. Alternatively modes 80 of climate variability which act on multi-decadal timescales, such as the Atlantic Multi-81 decadal Oscillation (Kerr, 2000; Deser & Phillips, 2021) and the Pacific Decadal Oscil-82 lation (Mantua & Hare, 2002), have an important imprint on Arctic sea ice. For exam-83 ple, variability emanating from the Pacific sector (Ding et al., 2018; Baxter et al., 2019) 84 or Atlantic sector (Meehl et al., 2018) has been suggested to have substantially contributed 85 to the rapid loss of Arctic sea ice during the 2000s (England et al., 2019). A number of 86 recent studies, using different methods including standard optimal detection method (Shen 87 et al., 2024), machine learning (Siew et al., 2024) and low-frequency component anal-88 ysis (Dörr et al., 2023), conclude that internal variability is at least as important as an-89 thropogenic forcing, perhaps more, for explaining the steep decline in that period. Need-90 less to say, internal variability can damp sea ice loss trends as well as strengthen them. 91 For instance, Yeager et al. (2015) correctly predicted a slowdown of winter Atlantic sec-92 tor sea ice loss for the past decade based on predictability from oceanic conditions linked 93 to the North Atlantic Oscillation. 94

In fact, it has been found in climate model simulations that internal climate vari-95 ability can totally counteract the forced loss of Arctic sea ice, resulting in periods of sim-96 ulated sea ice growth under increasing anthropogenic emissions. Kay et al. (2011) were 97 among the first to demonstrate, in a single climate model, that positive trends in Arc-98 tic sea ice extent on multi-decadal timescales were possible until the middle of this century. They found, using a limited ensemble size of six members, that two members ex-100 hibited statistically insignificant trends in September for the period 1979-2005 due to a 101 cancellation between the forced response and internal climate variability. Motivated by 102 the as-of-then brief pause in September Arctic sea ice loss for the period 2007-2013, Swart 103 et al. (2015) analyzed the CMIP5 archive and showed that seven-year pauses occurred 104 frequently in model simulations, and concluded that such episodes are an expected fea-105 ture of Arctic sea ice trajectory, even in a high emissions scenario. This study also demon-106 strated that pauses in sea ice loss on multi-decadal timescales remain plausible, and rel-107 atively frequent, over the coming century under a medium- or low-emissions scenario. 108 Looking back from the vantage point of 2025, the model-based studies of Kay et al. (2011) 109 and Swart et al. (2015) now appear remarkably prescient with regards to the plausibil-110 ity of a sustained slowdown in Arctic sea ice loss. 111

In this paper we document the recent observed multi-decadal pause in Arctic sea ice loss and address the following questions:

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1. How extensive and robust is this pause in Arctic sea ice loss?

- 2. Are comprehensive climate models able to capture this observed phenomenon, and
   if so how likely is it suggested to be?
- 3. How long could this observed pause plausibly persist for?
- 4. What is the role of anthropogenic forcing versus internal climate variability in contributing to the slow rate of sea ice loss?

#### <sup>120</sup> 2 Data and Methods

To investigate the evolution in Arctic sea ice cover, we utilise both the NSIDC (Fetterer 121 et al., 2017) and the OSISAF (OSI-420, 2023) sea ice indices. Both of these are contin-122 ually updated data records of the Arctic sea area and extent, for the period 1979 - present, 123 derived from satellite measurements. We note that there are known systematic differ-124 ences in the mean state between the two products (Meier & Stewart, 2019) but their inter-125 annual variations and multi-decadal trends have strong similarities (Figures 1a,b and S1). 126 For understanding changes in the simulated Arctic sea ice volume we utilise the Pan-Arctic 127 Ice Ocean Modeling and Assimilation System (PIOMAS) product (Schweiger et al., 2011). 128

To investigate the frequency and length of pauses in Arctic sea ice loss in compre-129 hensive climate model simulations, we here analyze all available large ensemble simula-130 tions from the CMIP5 and CMIP6 archive. Any model with at least ten members is used, 131 as summarised in Table S1. For the CMIP5 models, we use historical simulations which 132 terminate at the year 2005, followed by the Scenario MIP simulations with a range of Rep-133 resentative Concentration Pathways (RCPs). For the CMIP6 models, we use historical 134 simulations up to the year 2014, followed by ScenarioMIP simulations with a range of 135 Shared Socioeconomic Pathways (SSPs). 136

The main approach for analysing simulated changes in Arctic sea ice cover is to com-137 pute the linear trend for the twenty-year period 2005-2024 for each individual member 138 available for each model, as motivated by the observed changes (Section 3.1). This gives 139 a range between 10 and 100 members to examine the spread of simulated trends for each 140 model and scenario. To check for robustness by looking over a large sample size, we also 141 expand the overall time period by ten years each side (1995-2034), or shorter if the en-142 semble mean has transitioned to ice-free conditions before 2035, and then calculate all 143 of the possible 20-year trends during this period; this, however, does not substantially 144 alter the results. The main definition of slowdown used in this study is motivated by the 145 observed 2005-2024 September sea ice area trends (> -0.29 million km<sup>2</sup>/dec). We also 146 use an alternative definition – trends which are not statistically significant at the 95%147 confidence level - to ensure that this specific observed threshold does not overly influ-148 ence the results. This secondary definition contains information about the signal to noise 149 ratio, and so is complementary to the trend threshold definition. However, we find that 150 both definitions produce consistent results. 151

When we report multi-model averages, we do so by using a square-root weighting 152 scheme to take account of the number of members in each ensemble (models with more 153 members are weighted higher because the larger sample size will provide a more robust 154 estimate of the probability of a slowdown occurring) and the number of scenarios (mod-155 els with more scenarios are down-weighted because they are not independent of each other). 156 Doing this ensures that models with multiple scenarios are treated as if they have more 157 members of the same model scenario. The weighting for each model i and scenario j of 158 a given selection, where the number of members for each model for a given scenario is 159  $n_{ij}$  and the number of scenarios for each model is given by  $s_i$  is calculated as: 160

$$w_{ij} = \frac{\sqrt{a_{ij}}/\sqrt{s_i}}{\sum_{i,j}[\sqrt{a_{ij}}/\sqrt{s_i}]} \tag{1}$$

<sup>161</sup> However, we emphasize that this weighting scheme does not substantially alter the con-<sup>162</sup> clusions compared to if all members were weighted equally (not shown).

#### 163 **3 Results**

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#### 3.1 A robust and sustained pause in Arctic sea ice loss

We first investigate the recent observed trends in Arctic sea ice cover, focusing on 165 the annual minimum, September. The trend of September Arctic sea ice area for the most 166 recent two decades 2005-2024 is -0.30 and -0.29 million km<sup>2</sup> per decade according to the 167 NSIDC and OSISAF sea ice indices respectively (Fig. 1a,b). The key point, we empha-168 size, is that these trends are not statistically significantly different from zero at a 95%169 confidence level. This is also seen in Figure 1c,d where 20-year trends are plotted ver-170 sus the end year: note how trends ending in 2024 retreat inside the uncertainty envelope. 171 According to the OSISAF record, the 2005-2024 trend is the slowest rate of sea ice area 172 loss over any 20-year period since the start of the satellite record. For both datasets, this 173 insignificant trend is approximately four-times smaller than the peak 20-year sea ice loss 174 trend recorded (1993-2012). These results are robust to the choice of sea ice area or sea 175 ice extent (Fig. S1). The slowdown in September sea ice loss mainly occurs in the Pa-176 cific and Eurasian sector, from the Beaufort Sea westward to the Barents Sea (not shown). 177

While sea ice loss in September is of particular interest because that month is the annual minimum, the current pause in Arctic sea ice loss is seen in every single month throughout the year, as shown in Figure 1e,f). This suggests that the underlying mechanism(s) must explain not just the summer trends (R. Zhang, 2015; Francis & Wu, 2020) or winter trends (Yeager et al., 2015), but sea ice trends throughout the entire year.

The same picture - indicating a severe slowdown in Arctic sea ice loss - also emerges 183 when considering sea ice volume. The loss of Arctic sea ice volume has stalled for at least 184 the past fifteen years (Fig. S2). For the period 2010-2024, the simulated annual mean 185 Arctic sea ice volume has an approximately flat trend, decreasing by only 0.4 million km<sup>3</sup> 186 per decade, a value that is 7-times smaller than the long-term simulated loss for the pe-187 riod 1979-2024 of 2.9 million km<sup>3</sup> per decade, and again is not statistically significant. 188 This result, which is most evident in the Barents Sea (Onarheim et al., 2024), is consis-189 tent with a recent analysis suggesting a net build-up of sea ice volume since 2007 due 190 to a decrease in ice export from the Arctic, in addition to the thinner ice cover exhibit-191 ing higher growth rates (J. Zhang, 2021). 192

Given the strong observational evidence for a sustained and pervasive pause in Arctic sea ice loss over the recent 15-20 years, highly robust to the choice of sea ice metric, observational product, and season, we are led to ask: is such a pause unexpected? To answer that question we turn to analyzing comprehensive climate model simulations. We seek to determine if they are able to capture pauses such as the observed one and, if so, to establish if such pauses are exceedingly rare or relatively frequent events.

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#### 3.2 Comprehensive climate models suggest 20-year pauses are not rare

To understand whether comprehensive climate models can simulate a multi-decadal 200 pause of Arctic sea ice loss, we search through the CMIP5 and CMIP6 large ensemble 201 archive to identify members which exhibit ice loss pauses. Consistent with previous stud-202 ies (Kay et al., 2011; Swart et al., 2015; Lee & Liu, 2023), we find that nearly all mod-203 els are able to simulate reductions in September Arctic sea ice area smaller than observed 204 during the period 2005-2024. The two models which do not feature any such trends, UKESM1-205 0-LL and CanESM5-1, are both models with large climate sensitivities (Meehl et al., 2020), 206 for which overly strong anthropogenically-forced sea ice loss does not allow for pauses 207 such as the observed one. 208

Figure 2a shows the percentage of members with sea ice loss smaller than observed. The main result here is that the multi-model average suggests an approximately 20% chance of this pause in Arctic sea ice loss (Fig. 2a, column 1). However, we note a large spread



Figure 1. (a,b) Observed sea ice area  $[10^6 \text{ km}^2]$  1979-2024, (c,d) 20 year-trends of September sea ice area  $[10^6 \text{ km}^2/\text{decade}]$  with varying end year from 1998 to 2024, in which the red shaded envelope shows the bounds inside which a linear trend is not statistically significant according to a t-test at 95% confidence and (e,f) the 20 year-trends of sea ice area with varying end years but for each month of the year. The left column (a,c,e) shows the NSIDC sea index (Fetterer et al., 2017) and the right column (b,d,f) shows the OSISAF sea ice index (OSI-420, 2023).

across the CMIP5 and CMIP6 models, with the probability of a smaller-than-observed
2005-2024 trend varying from 0% and approximately 50%. Interestingly, the spread across
models for a given scenario is much larger than the spread across scenarios for a given
model. This is perhaps unsurprising because the scenarios diverge from each other later
than the 2020s (Notz & SIMIP Community, 2020), especially for the case of the CMIP6
forcing.

The central estimated value of approximately 20% doesn't change substantially if 218 models are selected following the criteria from Notz and SIMIP Community (2020), or 219 220 using models lying in the 66% range and 5-95% range estimates of the climate sensitivity (Sherwood et al., 2020), or using models according to their ability to simulate clima-221 tological sea ice conditions for the period 1979-1998 (Fig. 2a, column 2-4). Nor is the 222 central estimate substantially impacted if we assess the probability of a non-statistically 223 significant trend (Figure S3): this value is only slightly higher at approximately 25%. 224 Therefore from the multi-model perspective, what we have observed in the Arctic over 225 the past two decades is not a rare event, but rather one that should be expected to hap-226 pen with reasonable frequency. This result is insensitive to how models are sub-selected, 227 or to the metric of interest. This then raises the question of whether this pause in Arc-228 tic sea ice loss could continue and for how long. 229

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# **3.3** The observed pause in sea ice loss could foreseeably continue for another decade

To investigate how much longer this current pause is likely to last into the future, 232 we now examine those large ensemble members which do exhibit muted sea ice loss in 233 the period from 2005-2024 (Fig. 2a), and quantify how long the simulated slowdowns 234 persist in the future (Fig. 2b). In essence this produces the conditional probability es-235 timate of the 20-year pause extending further for each model and scenario. The multi-236 model average suggests that pauses in September sea ice loss for the period 2005-2024 237 have a 1 in 2 chance of persisting for a further five years, and a 1 in 3 chance of persist-238 ing for a further ten years (note, however, the considerable intermodel spread of  $\pm 25\%$ ). 239 On average, higher emissions scenarios tend to show slightly lower probabilities of sus-240 taining the muted pace of Arctic sea ice loss in the future, although the impact of this 241 is subtle and not consistent for every model. 242

On average, the sea ice area in ensemble members which simulate pauses in sea ice loss for the period 2005-2024 is 0.5 million km<sup>2</sup> larger in 2025 than in ensemble members from the same models and scenarios in which there is no pause (Fig. 2c). This source of predictability decays within a decade, and after that the September sea ice area in ensemble members with and without pauses are indistinguishable from each other.

It is important to highlight that to produce these estimates, we have limited the 248 model selection to only those models with at least five members which feature sea ice 249 pauses in 2004-2025, because to compute these probabilities in a meaningful way requires 250 that the remaining ensemble size is large enough. This may however, bias the results to-251 wards models with more ensemble members, which are more likely to include more mem-252 bers with slowdowns due to better sampling, and towards models which simulate slow-253 downs more frequently. The multi-model average of the probability of a sea ice loss pause 254 in this smaller subset of members is 28%, which is higher than the 20% estimated from 255 all models. If the observed slowdown is an inherently infrequent and rare event, then this 256 approach would overestimate the probability of it continuing, and underestimate how 257 anomalously high the Arctic sea ice cover is relative to the forced trend. However, of the 258 models which can reproduce the observed trends, over two-thirds of the available mod-259 els are included in this estimate. Our results are broadly consistent with the findings of 260 Swart et al. (2015), which showed that multi-decadal pauses longer than 20-years were 261 possible in the late 21st century under a medium emissions scenario. 262



Figure 2. (a) The percentage of ensemble members [%] for each ensemble that have 2005-2024 September sea ice area loss trends less than the observed value. The uncertainty estimate is calculated by monte carlo simulation with replacement. All simulations are shown on the left, with different selection criteria (that outlined in Notz and SIMIP Community (2020), the 5-95% and 66% range of ECS (Sherwood et al., 2020), and the climatological sea ice area applied on the right. (b) The conditional probability across each ensemble for the trends starting in 2005 to continue to be above -0.29 10<sup>6</sup> km<sup>2</sup>/decade for a given end year. (c) The ensemble-mean difference in September sea ice area [10<sup>6</sup> km<sup>2</sup>] between ensemble members with and without ice loss pauses over 2005-2024 for the period 2025-2050. For panels (b) and (c) only models and scenarios with at least five members with 2005-2024 trends above observed were included, with the black line shows the weighted average according to Equation 1. In all panels, each colour represents a model and each symbol represents a different forcing scenario.

Taken together, the wealth of available CMIP5 and CMIP6 simulations suggest it is possible, perhaps even likely, that the present slowdown in sea ice decline may continue for a further 5-10 years. If that were the case it may then imply the occurrence of an early ice-free Arctic is less likely than raw model output would suggest (Jahn et al., 2016; Arthun et al., 2020; Wang et al., 2021; England & Polvani, 2023; Jahn et al., 2024).

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#### 3.4 Climate models suggest an important role for climate variability

Whether the present slowdown persists in the future or not, one final question re-269 mains to be answered: is the recent pause a response to anthropogenic forcings alone, 270 or is there an important role for internal climate variability? When attempting to iso-271 late the forced component of any observed trend from internal variability, it is impor-272 tant to keep in mind that - assuming the model simulations faithfully capture a plau-273 sible reality - the observations are expected to have the same broad features as individ-274 ual ensemble members, i.e. that they are a combination of a forced trend plus one par-275 ticular realization of internal variability (although one doesn't expect the observations 276 to precisely match any one member). Given the well-established importance of internal 277 climate variability in Arctic sea ice trends (Kay et al., 2011; Swart et al., 2015; England 278 et al., 2019; Dörr et al., 2023), we next assess whether a change in the forced response 279 could also be substantially contributing to the observed slowdown in ice loss. 280

First, we show the forced September sea ice loss for the period 2005-2024 as esti-281 mated by the linear trend of the ensemble mean for each model and scenario (Fig. 3, hor-282 izontal axis). We find there are only two model/scenario combinations (GFDL-ESM2M 283 RCP8.5 and CESM2 SSP3-7.0) for which the forced trend is estimated to entirely ex-284 plain observed trends in ice loss (shown in Fig. 3, as grey vertical lines) with minimal 285 role for internal variability. Over 85% of the models we analyse here have a larger forced 286 sea ice loss for this period than observed (as they lie to the left of the observed trends), 287 implying that internal variability has acted to reduce the pace of ice loss. 288

Second, we ask: is there evidence that the forced response itself is slowing down 289 relative to the previous two decades? On the vertical axis of Figure 3, therefore, we plot 290 the ratio of the 2005-2024 forced trend to that of the preceding twenty years, 1986-2005, 291 for each model and scenario combination: a ratio of 1.0 indicates no change in the pace 292 of ice loss, > 1 indicates an acceleration and < 1 indicates a deceleration. Again we find 293 that only GFDL-ESM2M RCP8.5 and CESM2 SSP3-7.0 suggest that the reduction in 294 the forced trends accounts entirely for the observed slowdown. While the results from 295 all the other models agree that this observed pause is not entirely a forced response, the 296 remaining models are relatively evenly split (Fig. 3): roughly half the models suggest that 297 the forced sea ice loss trend has modestly decelerated over the past two decades relative 298 to the prior two decades, and roughly half suggest it has modestly accelerated (this is 200 especially clear if we disregard the models which are unable to capture the observed re-300 cent trends). 301

In summary then: while the modelling evidence is uncertain as to whether anthropogenic forcings - even in part - for the recent slowdown in Arctic sea ice loss, it is very likely that internal climate variability is contributing to the slowdown in an important way.

#### **306 4** Conclusion and Discussion

It is perhaps surprising that while global temperatures have risen rapidly, reaching record levels in the last few years, Arctic sea ice cover has shown no statistically significant decline over the past two decades. Nonetheless, analyzing two observational datasets and thousands of simulations from the CMIP5 and CMIP6 archives, we have established the following facts, which address the four questions raised in the introduction:



Figure 3. The ensemble mean trend in September sea ice area for the period 2005-2024 for each model and scenario (horizontal axis) versus the ratio of ensemble mean trends in September sea ice area for the periods 2005-2024 : 1986-2005 (vertical axis) with the black dotted line indicating a ratio of 1.0. Observational estimates of the 2005-2024 trend and the ratio of 2005-2024 : 1986-2005 trends are shown as grey lines (dashed line for NSIDC, solid line for OSISAF). Note that this does not imply that the observed trends are estimates of the forced response in the real climate system.

- 1. The pervasive slowdown of Arctic sea ice loss is robust across the choice of definitions, observational dataset, and season.
- 2. This observed pause in ice loss is simulated relatively frequently (with a 20% chance) 314 in climate models, and is thus to be expected even under high emission scenar-315 ios. 316
- 3. If model simulations are accurate, the recent pause may plausibly continue for an 317 additional five to ten years 318
  - 4. Nearly all models analysed suggest an important role for internal climate variability in slowing the anthropogenically-forced sea ice loss.

We now return to the question of the contribution of human influence versus in-321 ternal climate variability. If the slowdown is in fact a predominantly anthropogenically 322 forced episode, our results suggest that there must be either some shared missing forc-323 ing or common model deficiency in response to the standard forcing among the major-324 ity of the models. While the latter part is difficult to assess, one culprit for a missing forc-325 ing could be the increase in boreal forest fires, not incorporated in standard scenarios. 326 The recent study of Blanchard-Wrigglesworth et al. (2025) shows that incorporating re-327 cent biomass burning emissions into the CESM2 model leads to a rapid recovery of Septem-328 ber sea ice cover during the period of interest due to increased reflection in the North-329 ern Hemisphere of incoming shortwave radiation arising from the cloud response and aerosol 330 cloud interactions. However, due to specifics of the simulation of polar clouds in CESM2 331 (DeRepentigny et al., 2022; Zhu et al., 2022; Davis & Medeiros, 2024; England & Feldl, 332 2024), and a seemingly opposite result from similar experiments with a different climate 333 model (Zhong et al., 2024), further experiments with a wider range of models are needed 334 to understand the role changes in biomass burning have had on observed Arctic sea ice 335 trends. 336

Going forward, how can we use what we have learned about the recent pause in 337 Arctic sea ice loss? Firstly, if internal variability has played an important role then this 338 could provide a source of future predictability of Arctic climate change in the same man-339 ner as Yeager et al. (2015). And second, this period could be used as an out of sample 340 test in future climate model evaluation – similar to the early and middle periods of the 341 20th century (Flynn et al., 2023; Bianco et al., 2024; Chen & Dai, 2024). However, over-342 all this study is a reminder that we should be humble about multi-decadal predictions 343 of the climate system, especially in highly variable regions such as the Arctic. Standing in 2007 or 2012 after having experienced another year of record loss and listening to as-345 sessments that climate models are flawed in their ability to reproduce the rapid loss of 346 Arctic sea ice (Stroeve et al., 2007), it would take a rather brave person to have predicted 347 that a sustained slowdown in ice loss was around the corner, although, as we have shown, 348 and many have found before (Kay et al., 2011; Swart et al., 2015; R. Zhang, 2015), this 349 is entirely consistent with what climate models simulate. 350

#### **Open Research Section** 351

All CMIP5 and CMIP6 data analysed in this study is publicly available to down-352 load from the Earth System Federation Grid at https://aims2.llnl.gov/search. 353

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#### 360 References

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- 361Arthun, M., Onarheim, I., Dörr, J., & Eldevik, T.(2020).The seasonal and362regional transition to an ice-free Arctic.Geophysical Research Letters,36348 (e2020GL090825).doi: 10.1029/2020GL0909825
- Baxter, I., Ding, Q., Schweiger, A., L'Heureux, M., Baxter, S., Wang, T., ... Lu, J.
   (2019). How tropical Pacific surface cooling contributed to accelerated sea ice melt from 2007 to 2012 as ice is thinned by anthropogenic forcing. *Journal of Climate*, 32, 8583-8602. doi: 10.1175/JCLI-D-18-0783.1
- Bianco, E., Blanchard-Wrigglesworth, E., Materia, S., Ruggieri, P., Iovino, D.,
  - & Masina, S. (2024). CMIP6 models underestimate Arctic sea ice loss during the early twentieth-century warming, despite simulating large low-
  - frequency sea ice variability. Journal of Climate, 37, 6305-6321. doi: 10.1175/JCLI-D-23-0647.1
- Blanchard-Wrigglesworth, E., DeRepentigny, P., & Frierson, D. (2025). Increasing
   boreal fires reduce future global warming and sea ice loss. *Preprint*. doi: 10
   .21203/rs.3.rs-5486231/v2
- Chen, X., & Dai, A. (2024). Quantifying contributions of external forcing and internal variability to Arctic warming during 1900-2021. *Earth's Future*, *12*(e2023EF003734). doi: 10.1029/2023EF003734
- Copernicus, E. (2024, April). *Climate indicators.* Retrieved 16/02/2025, from https://climate.copernicus.eu/climate-indicators
  - Davis, I., & Medeiros, B. (2024). Assessing CESM2 clouds and their response to climate change using cloud regimes. Journal of Climate, 37, 2965-2985. doi: 10 .1175/JCLI-D-23-0337.1
- DeRepentigny, P., Jahn, A., Holland, M., Kay, J., Fasullo, J., Lamarque, J., ... Barrett, A. (2022). Enhanced simulated early 21st century Arctic sea ice loss due to CMIP6 biomass burning emissions. *Science Advances*, 8(30). doi: 10.1126/sciadv.abo2405
  - Deser, C., & Phillips, A. (2021). Defining the internal component of Atlantic Multidecadal Variability in a changing climate. Geophysical Research Letters, 48 (e2021GL095023). doi: 10.1029/2021GL095023
  - Ding, Q., Schweiger, A., L'Heureux, M., Steig, E., Battisti, D., Johnson, N., ...
  - Baxter, I. (2018). Fingerprints of internal drivers of Arctic sea ice loss in observations and model simulations. *Nature Geoscience*, 12, 28-33. doi: 10.1038/s41561-018-0256-8
- Dörr, J., Bonan, D., Arthun, M., Svendsen, L., & Wills, R. (2023). Forced and inter nal components of observed Arctic sea-ice changes. *The Cryosphere*, 17, 4133 4153. doi: 10.5194/tc-17-4133-2023
- England, M. (2021). Are multi-decadal fluctuations in arctic and antarctic surface
   temperatures a forced response to anthropogenic emissions or part of internal
   climate variability? *Geophysical Research Letters*, 48 (e2020GL090631). doi:
   10.1029/2020GL090631
- England, M., Eisenman, I., Lutsko, N., & Wagner, T. (2021). The recent emergence
   of Arctic Amplification. *Geophysical Research Letters*, 48 (e2021GL094086).
   doi: 10.1029/2021GL094086
- England, M., & Feldl, N. (2024). Robust polar amplification in ice-free climates
   relies on ocean heat transport and cloud radiative effects. Journal of Climate,
   37, 2179-2197. doi: 10.1175/JCLI-D-23-0151.1
- England, M., Jahn, A., & Polvani, L. (2019). Nonuniform contribution of internal
   variability to recent Arctic sea ice loss. Journal of Climate, 32, 4039-4053. doi:
   10.1175/JCLI-D-18-0864.1
- England, M., & Polvani, L. (2023). The Montreal Protocol is delaying the occurrence of the first ice-free Arctic summer. *PNAS*, 120 (e2211432120). doi: 10
  .1073/pnas.2211432120

Fetterer, F., Knowles, K., Meier, W., Savoie, M., & Windnagel, A. (2017). Sea ice 414 index. (q02135, version 3). [dataset]. National Snow and Ice Data Center. doi: 415 10.7265/N5K072F8 416 Flynn, C., Huusko, L., Modak, A., & Mauritsen, T. (2023).Strong aerosol cool-417 ing alone does not explain cold-biased mid-century temperatures in CMIP6 418 models. Atmospheric Chemistry and Physics, 23, 15,121-15,133. doi: 419 10.5194/acp-23-15121-2023 420 Francis, J., & Wu, B. (2020). Why has no new record-minimum Arctic sea-ice extent 421 occurred since september 2012? Environmental Research Letters, 15(114034). 422 doi: 10.1088/1748-9326/abc047 423 Fyfe, J., von Salzen, K., Gillett, N., Arora, V., Flato, G., & McConnell, J. (2013).424 One hundred years of Arctic surface temperature variation due to anthro-425 pogenic influence. Scientific Reports, 3(2645). doi: 10.1038/srep02645 426 Gagne, M., Fyfe, J., Gillett, N., Polyakov, I., & Flato, G. (2017).Aerosol-driven 427 increase in Arctic sea ice over the middle of the twentieth century. Geophysical 428 Research Letters, 44, 7338-7346. doi: 10.1002/2016GL071941 429 Hansen, J., Kharecha, P., Sato, M., Tselioudis, G., Kelly, J., Bauer, S., ... Pokela, 430 Α. (2025).Global warming has accelerated: Are the United Nations and 431 the public well-informed? Environment: Science and Policy for Sustainable 432 Development, 67, 6-44. doi: 10.1080/00139157.2025.2434494 433 IPCC. (2021).Climate change 2021: The physical science basis. Contribution of 434 working group I to the sixth assessment report of the Intergovernmental Panel 435 on Climate Change. In V. Masson-Delmotte et al. (Eds.), (p. 3-32). Cambridge 436 University Press. 437 (2024).Jahn, A., Holland, M., & Kay, J. Projections of an ice-free Arctic Ocean. 438 Nature Reviews Earth and Environment, 5, 167-176. doi: 10.1038/s43017-023 439 -00515-9440 Jahn, A., Kay, J., Holland, M., & Hall, D. (2016). How predictable is the timing of a 441 summer ice-free Arctic? Geophysical Research Letters, 43, 9113-9120. doi: 10 442 .1002/2016GL070067 443 Kay, J., Holland, M., & Jahn, A. (2011). Inter-annual to multi-decadal Arctic sea ice 444 extent trends in a warming world. Geophysical Research Letters, 38(L15708). 445 doi: 10.1029/2011GL048008 446 Kerr, R. (2000). A North Atlantic climate pacemaker for the centuries. Science. 447 288, 1984-1985. doi: 10.1126/science.288.5473.1984 448 Kwok, R. (2018). Arctic sea ice thickness, volume, and multiyear ice coverage: losses 449 and coupled variability (1958–2018). Environmental Research Letters, 13, 450 105005. doi: 10.1088/1748-9326/aae3ec 451 Lee, Y., & Liu, W. (2023).The weakened Atlantic Meridional Overturning Cir-452 culation diminishes recent Arctic sea ice loss. Geophysical Research Letters, 453 50(e2023GL105929). doi: 10.1029/2023GL105929 454 Anthropogenic aerosol forcing under Lund, M., Myhre, G., & Samset, B. (2019).455 the Shared Socioeconomic Pathways. Atmospheric Chemistry and Physics, 19, 456 13827-13839. doi: 10.5194/acp-19-13827-2019 457 Manshausen, P., Watson-Parris, D., Christensen, M., Jalkanen, J., & Stier, P. 458 (2022).Invisible ship tracks show large cloud sensitivity to aerosol. Nature. 459 610, 101–106. doi: 10.1038/s41586-022-05122-0 460 Mantua, N., & Hare, S. (2002). The Pacific Decadal Oscillation. Journal of Oceanog-461 raphy, 58, 35-44. doi: 10.1023/A:1015820616384 462 Maslowski, W., Kinney, J., Higgins, M., & Roberts, A. (2012). The future of Arctic 463 sea ice. Annual Review of Earth and Planetary Sciences, 40, 625-654. doi: 10 464 .1146/annurev-earth-042711-105345 465 Meehl, G., Chung, C., Arblaster, J., Holland, M., & Bitz, C. (2018).Tropical 466 decadal variability and the rate of Arctic sea ice decrease. Geophysical Re-467 search Letters, 45, 11,326-11,333. doi: 10.1029/2018GL079989 468

469	Meehl, G., Senior, C., Eyring, V., Lamarque, J., Stouffer, R., Taylor, K., & Schlund,
470	M. (2020). Context for interpreting equilibrium climate sensitivity and tran-
471	sient climate response from the CMIP6 Earth system models. Science Ad-
472	vances, 6 (eaba1981). doi: 10.1126/sciadv.aba1981
473	Meier, W., & Stewart, J. (2019). Assessing uncertainties in sea ice extent climate
474	indicators. Environmental Research Letters, 14(035005). doi: 10.1088/1748
475	-9326/aaf52c
476	Merchant, C., Allan, R., & Embury, O. (2025). Quantifying the acceleration of mul-
477	tidecadal global sea surface warming driven by Earth's energy imbalance. En-
478	vironmental Research Letters, 20(024037). doi: 10.1088/1748-9326/adaa8a
479	Nafaji, M., Zwiers, F., & Gillett, N. (2015). Attribution of Arctic temperature
480	change to greenhouse-gas and aerosol influences. Nature Climate Change, 5.
481	246-249. doi: 10.1038/nclimate2524
482	Notz, D., & SIMIP Community. (2020). Arctic sea ice in CMIP6. Geophysical Re-
483	search Letters, 47(e2019GL086749). doi: 10.1029/2019GL086749
484	Notz, D., & Stroeve, J. (2016). Observed Arctic sea-ice loss directly follows anthro-
485	pogenic CO2 emission. <i>Science</i> , 354, 747-750, doi: 10.1126/science.aag2345
196	Onarheim I Arthun M Teigen S Eik K & Steele M (2024) Recent
400	thickening of the Barents Sea ice cover Geonbusical Research Letters
401	51 (e2024GL108225) doi: 10 1029/2024GL108225
400	OSI-420 (2023) Osi saf sea ice inder 1978-onwards version 2.2 [dataset] EU-
409	METSAT Ocean and Sea Ice Satellite Application Facility, doi: 10.15770/EUM
490	SAF OSI 0022
491	Parkinson C & Comiso I (2013) On the 2012 record low Arctic sea ice cover:
492	Combined impact of preconditioning and an August storm <i>Ceonbusical Re-</i>
493	search Letters 10 1356-1361 doi: 10 1002/grl 50349
494	Bantanan M. Karpachko A. Lipponan A. Huyarinan O. Buostaanaja K.
495	Vibra T & Laaksonen $\Lambda$ (2022) The Arctic has warmed nearly four times
496	faster than the globe since 1070 Communications Earth and Environment
497	3(168) doi: 10.1038/s/3207-022-00/08-3
498	Samset B. Zhou, C. Fuglestvedt I. Lund M. Marotzka, I. & Zelinka, M. (2023)
499	Steady global surface warming from 1073 to 2022 but increased warming
500	rate after 1990 Communications Earth and Environment /(400) doi:
501	$10 \ 1038 \ s 43247.023.01061.4$
502	Schweiger A Lindsey R Zhang I Steele M Stern H & Kwok R (2011) Un-
503	certainty in modeled Arctic sea ice volume Journal of Geophysical Research
504	116(C00D06) doi: 10.1029/2011.IC007084
505	Shen Z. Duan A. Zhou W. Peng V. & Li I. (2024). Reconciling roles of ex-
500	ternal forcing and internal variability in Arctic sea ice change on different time
507	scales Lowrnal of Climate 37 3577-3591 doi: 10.1175/JCLI-D-23-0280.1
500	Sherwood S Webb M Annan I Armour K Forster P Hargreaves I
509	Zelinka M (2020) An assessment of Earth's climate sensitivity using mul-
510	tiple lines of evidence <i>Review of Geophysics</i> 58(e2019BG000678) doi:
511	10 1029/2019BG000678
512	Siew P. Wu V. Ting M. Zheng C. Ding O. & Seager B. (2024) Signif-
515	icant contribution of internal variability to recent Barents-Kara sea ice
514	loss in winter <u>Communications Earth and Environment</u> 5(411) doi:
515	10 1038/s43247_024_01582_6
510	Stroeve I Holland M Mejer W Scambos T & Sarrozo M (2007) Arctic
517	sea ice decline: faster than forecast. <i>Geophysical Research Letters</i> 9/(L00501)
510	doi: 10.1029/2007GL029703
213	Stroeve I & Notz D (2018) Changing state of Arctic con ico across all concord
520	Environmental Research Letters 12(103001) doi: 10.1082/1748.0296/204056
521	Stroeve I Serreze M Drobot S Coerboad S Holland M Madanika I
522	Scombos T (2011) Arctia con ico extent plummete in 2007 Eco 20 12 14
523	Scambos, 1. (2011). Arctic sea ice extent plummets in 2007. $Eos, \delta 9, 13-14$ .

524	doi: 10.1029/2008EO020001
525	Swart, N., Fyfe, J., Hawkins, E., Kay, J., & Jahn, A. (2015). Influence of internal
526	variability on arctic sea-ice trends. Nature Climate Change, 5, 86-89. doi: 10
527	.1038/nclimate2483
528	Szopa, S., Balkanski, Y., Schulz, M., Bekki, S., Cugnet, D., Fortems-Cheiney, A.,
529	Dufresne, J. (2013). Aerosol and ozone changes as forcing for climate
530	evolution between 1850 and 2100. Climate Dynamics, 40, 2223-2250. doi:
531	10.1007/s00382-012-1408-y
532	Wadhams, P. (2016). A farewell to ice. Penguin Publishing.
533	Walsh, J., Fetterer, F., Scott, J., & Chapman, W. (2017). A database for depicting
534	Arctic sea ice variations back to 1850. Geographical Review, 107, 89-107. doi:
535	10.1111/j.1931-0846.2016.12195.x
536	Wang, B., Zhou, X., Ding, Q., & Liu, J. (2021). Increasing confidence in project-
537	ing the Arctic ice-free year with emergent constraints. Environmental Research
538	Letters, $16(094016)$ . doi: $10.1088/1748-9326/ac0b17$
539	Yeager, S., Karspeck, A., & Danabasoglu, G. (2015). Predicted slowdown in the rate
540	of Atlantic sea ice loss. Geophysical Research Letters, 42, 10,704-10,713. doi:
541	$10.1002/2015 { m GL}065364$
542	Yoshioka, M., Grosvenor, D., Booth, B., Morice, C., & Carslaw, K. (2024). Warming
543	effects of reduced sulfur emissions from shipping. Atmospheric Chemistry and
544	<i>Physics</i> , $24$ , 13681–13692. doi: 10.5194/acp-24-13681-2024
545	Zhang, J. (2021). Recent slowdown in the decline of Arctic sea ice volume under
546	increasingly warm atmospheric and oceanic conditions. Geophysical Research
547	<i>Letters</i> , 48(e2021GL094780). doi: 10.1029/2021GL094780
548	Zhang, J., Lindsay, R., Schweiger, A., & Steele, M. (2013). The impact of an intense
549	summer cyclone on 2012 Arctic sea ice retreat. Geophysical Research Letters,
550	40,720-726. doi: 10.1002/gr1.50190
551	Zhang, R. (2015). Mechanisms for low-frequency variability of summer Arctic sea ice
552	extent. PNAS, 112, 4570-4575. doi: 10.1073/pnas.1422296112
553	Zhong, Q., Shutgens, N., Veraverbeke, S., & van der Werf, G. (2024). Increasing
554	aerosol emissions from boreal biomass burning exacerbate Arctic warming. Na-
555	ture Change, 14, 1275-1281. doi: 10.1038/s41558-024-02176-y
556	Linu, J., Otto-Biesner, B., Brady, E., Gettelman, A., Bacmeister, J., Neale, R.,
557	ray, J. (2022). LGW paleoclimate constraints inform cloud parameteriza-
558	tions and equilibrium chinate sensitivity in CESM2. Journal of Advances in Modeling Forth Systems, 17(20021MS002776), doi: 10.1020/2021MS002776
559	<i>Modeling Earth Systems</i> , 14 (e2021MS002776). doi: 10.1029/2021MS002776

#### Surprising, but not unexpected, multi-decadal pause in 1 Arctic sea ice loss 2

M. R. England<sup>1</sup>, L. M. Polvani<sup>2,3</sup>, J. Screen<sup>1</sup>, and A. C. Chan<sup>1</sup>

<sup>1</sup>Department of Mathematics and Statistics, University of Exeter, UK <sup>2</sup>Lamont-Doherty Earth Observatory, Columbia University, New York, USA <sup>3</sup>Department of Applied Physics and Applied Mathematics, Columbia University, New York, USA

## Key Points:

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8	•	The loss of Arctic sea ice cover has undergone a pronounced slowdown over the
9		past two decades, across all months of the year.
10	•	Rather than being an unexpected rare event, comprehensive climate models from
11		CMIP5 and CMIP6 simulate such pauses relatively frequently.
12	•	According to these climate model simulations, this pause in the loss of Arctic sea
13		ice could plausibly continue for the next 5-10 years.

Corresponding author: Mark Ross England, m.england2@exeter.ac.uk

#### 14 Abstract

Over the past two decades, Arctic sea ice loss has slowed considerably, with no statis-15 tically significant decline in September sea ice area since 2005. This pause is robust across 16 observational datasets, metrics, and seasons. Large-ensemble CMIP5 and CMIP6 sim-17 ulations reveal that such periods with no sea ice decline under increasing greenhouse gas 18 emissions are not unusual. Analysis of ensemble members that simulate analogues of the 19 observed pause indicates that the current slowdown could plausibly persist another five 20 to ten years. The modelling evidence suggests that internal variability has substantially 21 offset anthropogenically forced sea ice loss in recent decades, although possible contri-22 butions from changes in the forced response remain uncertain. Overall, this observed pause 23 in Arctic sea ice decline is consistent with simulated internal variability superimposed 24 on the long term trend according to the bulk of the climate modelling evidence. 25

#### <sup>26</sup> Plain Language Summary

Over the last 20 years, the decline of Arctic sea ice has slowed down substantially. 27 Climate models (from CMIP5 and CMIP6) show that pauses in sea ice loss across mul-28 tiple decades can happen, even as greenhouse gas emissions continue to rise. When we 29 compare the current slowdown to similar pauses in model simulations, we see that this 30 "hiatus" could plausibly continue for another five to ten years. Most of the evidence from 31 these climate models suggests that natural climate variations have played a large part 32 in slowing the human-driven loss of sea ice. However, it is not entirely certain whether 33 changes in the human influence on climate (the "forced response") have also contributed. 34 Overall, while it may sound surprising that Arctic sea ice loss has slowed down even as 35 global temperatures hit record highs, the climate modelling evidence suggests we should 36 expect periods like this to occur somewhat frequently. 37

#### 38 1 Introduction

The loss of Arctic sea ice over the past half century is one of the most clear and 39 well-known indicators of human-induced climate change (IPCC, 2021; Copernicus, 2024). 40 September sea ice area has nearly halved since the beginning of the satellite era in 1979 41 (Fetterer et al., 2017; Stroeve & Notz, 2018), and during the same period, estimated Arc-42 tic sea ice volume has decreased by over 10,000 km<sup>3</sup> (Kwok, 2018). Record-breaking sum-43 mer sea ice minimums in 2007 (Stroeve et al., 2011) and 2012 (Parkinson & Comiso, 2013; 44 J. Zhang et al., 2013) fuelled predictions, which with hindsight look overly alarmist, that 45 the Arctic would experience its first ice-free summer before 2020 (Maslowski et al., 2012; 46 Wadhams, 2016). Adding to this, the Arctic has been warming up to four times faster 47 than the global average (Rantanen et al., 2022). It has been further proposed that global 48 warming might be accelerating, culminating in record breaking warmth in recent years 49 (Samset et al., 2023; Hansen et al., 2025; Merchant et al., 2025). As Arctic sea ice cover 50 is strongly tied to global temperatures (Notz & Stroeve, 2016), there would be little ex-51 pectation of a multi-decadal *slowdown* in Arctic sea ice loss. And yet, as we will show, 52 such a slowdown has been occurring in the last two decades. 53

Recall that, over the past century, periods of increasing anthropogenic greenhouse 54 emissions without sustained sea ice loss - the mid-20th century (Walsh et al., 2017) - have 55 already occurred. From the 1940s to the 1970s Arctic sea ice cover expanded (Gagne et 56 al., 2017), with the largest increases in the Chukchi, East Siberian, Laptev, Kara and 57 Barents Seas. However, anthropogenic forcing in the mid-20th century was very differ-58 ent compared to the one of the past two decades. Industrial aerosol emissions from Eu-59 rope and North America contributed substantially to the positive multi-decadal trend 60 in Arctic sea ice area and associated Arctic cooling in the mid-20th century (Fyfe et al., 61 2013; Nafaji et al., 2015; Gagne et al., 2017; England et al., 2021); but these aerosol sources 62 are far smaller today (Szopa et al., 2013; Lund et al., 2019). However, when anthropogenic 63

aerosols are being discussed in the context of the recent past, it is with regards to the
phase out of aerosol emissions from ship tracks which have potentially contributed to enhanced global warming since 2020 (Manshausen et al., 2022; Yoshioka et al., 2024). In
fact, Yoshioka et al. (2024) find that the simulated warming response to these reduced
sulphur emissions is largest in the Arctic. So, the lessons of the past may not be a reliable guide for understanding current trends.

It is important to appreciate that the observed trend in Arctic sea ice cover over 70 a given period is composed of a contribution caused by anthropogenic emissions, denoted 71 72 the forced response, and a contribution from unforced fluctuations associated with internal climate variability (England et al., 2019; England, 2021; Dörr et al., 2023; Shen 73 et al., 2024). Anthropogenically-forced changes which may contribute to a reduction in 74 Arctic sea ice loss over the past two decades include a forced slowdown in the Atlantic 75 Meridional Overturning Circulation (Lee & Liu, 2023), and changes in the emissions from 76 biomass burning, both in the magnitude (Blanchard-Wrigglesworth et al., 2025), and the 77 variability (DeRepentigny et al., 2022). One would imagine, however, that the aforemen-78 tioned reduction of sulphur emission from shiptracks (Yoshioka et al., 2024) would lead 79 to an acceleration rather than a deceleration of sea ice loss since 2020. Alternatively modes 80 of climate variability which act on multi-decadal timescales, such as the Atlantic Multi-81 decadal Oscillation (Kerr, 2000; Deser & Phillips, 2021) and the Pacific Decadal Oscil-82 lation (Mantua & Hare, 2002), have an important imprint on Arctic sea ice. For exam-83 ple, variability emanating from the Pacific sector (Ding et al., 2018; Baxter et al., 2019) 84 or Atlantic sector (Meehl et al., 2018) has been suggested to have substantially contributed 85 to the rapid loss of Arctic sea ice during the 2000s (England et al., 2019). A number of 86 recent studies, using different methods including standard optimal detection method (Shen 87 et al., 2024), machine learning (Siew et al., 2024) and low-frequency component anal-88 ysis (Dörr et al., 2023), conclude that internal variability is at least as important as an-89 thropogenic forcing, perhaps more, for explaining the steep decline in that period. Need-90 less to say, internal variability can damp sea ice loss trends as well as strengthen them. 91 For instance, Yeager et al. (2015) correctly predicted a slowdown of winter Atlantic sec-92 tor sea ice loss for the past decade based on predictability from oceanic conditions linked 93 to the North Atlantic Oscillation. 94

In fact, it has been found in climate model simulations that internal climate vari-95 ability can totally counteract the forced loss of Arctic sea ice, resulting in periods of sim-96 ulated sea ice growth under increasing anthropogenic emissions. Kay et al. (2011) were 97 among the first to demonstrate, in a single climate model, that positive trends in Arc-98 tic sea ice extent on multi-decadal timescales were possible until the middle of this century. They found, using a limited ensemble size of six members, that two members ex-100 hibited statistically insignificant trends in September for the period 1979-2005 due to a 101 cancellation between the forced response and internal climate variability. Motivated by 102 the as-of-then brief pause in September Arctic sea ice loss for the period 2007-2013, Swart 103 et al. (2015) analyzed the CMIP5 archive and showed that seven-year pauses occurred 104 frequently in model simulations, and concluded that such episodes are an expected fea-105 ture of Arctic sea ice trajectory, even in a high emissions scenario. This study also demon-106 strated that pauses in sea ice loss on multi-decadal timescales remain plausible, and rel-107 atively frequent, over the coming century under a medium- or low-emissions scenario. 108 Looking back from the vantage point of 2025, the model-based studies of Kay et al. (2011) 109 and Swart et al. (2015) now appear remarkably prescient with regards to the plausibil-110 ity of a sustained slowdown in Arctic sea ice loss. 111

In this paper we document the recent observed multi-decadal pause in Arctic sea ice loss and address the following questions:

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1. How extensive and robust is this pause in Arctic sea ice loss?

- 2. Are comprehensive climate models able to capture this observed phenomenon, and
   if so how likely is it suggested to be?
- 3. How long could this observed pause plausibly persist for?
- 4. What is the role of anthropogenic forcing versus internal climate variability in contributing to the slow rate of sea ice loss?

#### <sup>120</sup> 2 Data and Methods

To investigate the evolution in Arctic sea ice cover, we utilise both the NSIDC (Fetterer 121 et al., 2017) and the OSISAF (OSI-420, 2023) sea ice indices. Both of these are contin-122 ually updated data records of the Arctic sea area and extent, for the period 1979 - present, 123 derived from satellite measurements. We note that there are known systematic differ-124 ences in the mean state between the two products (Meier & Stewart, 2019) but their inter-125 annual variations and multi-decadal trends have strong similarities (Figures 1a,b and S1). 126 For understanding changes in the simulated Arctic sea ice volume we utilise the Pan-Arctic 127 Ice Ocean Modeling and Assimilation System (PIOMAS) product (Schweiger et al., 2011). 128

To investigate the frequency and length of pauses in Arctic sea ice loss in compre-129 hensive climate model simulations, we here analyze all available large ensemble simula-130 tions from the CMIP5 and CMIP6 archive. Any model with at least ten members is used, 131 as summarised in Table S1. For the CMIP5 models, we use historical simulations which 132 terminate at the year 2005, followed by the Scenario MIP simulations with a range of Rep-133 resentative Concentration Pathways (RCPs). For the CMIP6 models, we use historical 134 simulations up to the year 2014, followed by ScenarioMIP simulations with a range of 135 Shared Socioeconomic Pathways (SSPs). 136

The main approach for analysing simulated changes in Arctic sea ice cover is to com-137 pute the linear trend for the twenty-year period 2005-2024 for each individual member 138 available for each model, as motivated by the observed changes (Section 3.1). This gives 139 a range between 10 and 100 members to examine the spread of simulated trends for each 140 model and scenario. To check for robustness by looking over a large sample size, we also 141 expand the overall time period by ten years each side (1995-2034), or shorter if the en-142 semble mean has transitioned to ice-free conditions before 2035, and then calculate all 143 of the possible 20-year trends during this period; this, however, does not substantially 144 alter the results. The main definition of slowdown used in this study is motivated by the 145 observed 2005-2024 September sea ice area trends (> -0.29 million km<sup>2</sup>/dec). We also 146 use an alternative definition – trends which are not statistically significant at the 95%147 confidence level - to ensure that this specific observed threshold does not overly influ-148 ence the results. This secondary definition contains information about the signal to noise 149 ratio, and so is complementary to the trend threshold definition. However, we find that 150 both definitions produce consistent results. 151

When we report multi-model averages, we do so by using a square-root weighting 152 scheme to take account of the number of members in each ensemble (models with more 153 members are weighted higher because the larger sample size will provide a more robust 154 estimate of the probability of a slowdown occurring) and the number of scenarios (mod-155 els with more scenarios are down-weighted because they are not independent of each other). 156 Doing this ensures that models with multiple scenarios are treated as if they have more 157 members of the same model scenario. The weighting for each model i and scenario j of 158 a given selection, where the number of members for each model for a given scenario is 159  $n_{ij}$  and the number of scenarios for each model is given by  $s_i$  is calculated as: 160

$$w_{ij} = \frac{\sqrt{a_{ij}}/\sqrt{s_i}}{\sum_{i,j}[\sqrt{a_{ij}}/\sqrt{s_i}]} \tag{1}$$

<sup>161</sup> However, we emphasize that this weighting scheme does not substantially alter the con-<sup>162</sup> clusions compared to if all members were weighted equally (not shown).

#### 163 **3 Results**

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#### 3.1 A robust and sustained pause in Arctic sea ice loss

We first investigate the recent observed trends in Arctic sea ice cover, focusing on 165 the annual minimum, September. The trend of September Arctic sea ice area for the most 166 recent two decades 2005-2024 is -0.30 and -0.29 million km<sup>2</sup> per decade according to the 167 NSIDC and OSISAF sea ice indices respectively (Fig. 1a,b). The key point, we empha-168 size, is that these trends are not statistically significantly different from zero at a 95%169 confidence level. This is also seen in Figure 1c,d where 20-year trends are plotted ver-170 sus the end year: note how trends ending in 2024 retreat inside the uncertainty envelope. 171 According to the OSISAF record, the 2005-2024 trend is the slowest rate of sea ice area 172 loss over any 20-year period since the start of the satellite record. For both datasets, this 173 insignificant trend is approximately four-times smaller than the peak 20-year sea ice loss 174 trend recorded (1993-2012). These results are robust to the choice of sea ice area or sea 175 ice extent (Fig. S1). The slowdown in September sea ice loss mainly occurs in the Pa-176 cific and Eurasian sector, from the Beaufort Sea westward to the Barents Sea (not shown). 177

While sea ice loss in September is of particular interest because that month is the annual minimum, the current pause in Arctic sea ice loss is seen in every single month throughout the year, as shown in Figure 1e,f). This suggests that the underlying mechanism(s) must explain not just the summer trends (R. Zhang, 2015; Francis & Wu, 2020) or winter trends (Yeager et al., 2015), but sea ice trends throughout the entire year.

The same picture - indicating a severe slowdown in Arctic sea ice loss - also emerges 183 when considering sea ice volume. The loss of Arctic sea ice volume has stalled for at least 184 the past fifteen years (Fig. S2). For the period 2010-2024, the simulated annual mean 185 Arctic sea ice volume has an approximately flat trend, decreasing by only 0.4 million km<sup>3</sup> 186 per decade, a value that is 7-times smaller than the long-term simulated loss for the pe-187 riod 1979-2024 of 2.9 million km<sup>3</sup> per decade, and again is not statistically significant. 188 This result, which is most evident in the Barents Sea (Onarheim et al., 2024), is consis-189 tent with a recent analysis suggesting a net build-up of sea ice volume since 2007 due 190 to a decrease in ice export from the Arctic, in addition to the thinner ice cover exhibit-191 ing higher growth rates (J. Zhang, 2021). 192

Given the strong observational evidence for a sustained and pervasive pause in Arctic sea ice loss over the recent 15-20 years, highly robust to the choice of sea ice metric, observational product, and season, we are led to ask: is such a pause unexpected? To answer that question we turn to analyzing comprehensive climate model simulations. We seek to determine if they are able to capture pauses such as the observed one and, if so, to establish if such pauses are exceedingly rare or relatively frequent events.

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#### 3.2 Comprehensive climate models suggest 20-year pauses are not rare

To understand whether comprehensive climate models can simulate a multi-decadal 200 pause of Arctic sea ice loss, we search through the CMIP5 and CMIP6 large ensemble 201 archive to identify members which exhibit ice loss pauses. Consistent with previous stud-202 ies (Kay et al., 2011; Swart et al., 2015; Lee & Liu, 2023), we find that nearly all mod-203 els are able to simulate reductions in September Arctic sea ice area smaller than observed 204 during the period 2005-2024. The two models which do not feature any such trends, UKESM1-205 0-LL and CanESM5-1, are both models with large climate sensitivities (Meehl et al., 2020), 206 for which overly strong anthropogenically-forced sea ice loss does not allow for pauses 207 such as the observed one. 208

Figure 2a shows the percentage of members with sea ice loss smaller than observed. The main result here is that the multi-model average suggests an approximately 20% chance of this pause in Arctic sea ice loss (Fig. 2a, column 1). However, we note a large spread



Figure 1. (a,b) Observed sea ice area  $[10^6 \text{ km}^2]$  1979-2024, (c,d) 20 year-trends of September sea ice area  $[10^6 \text{ km}^2/\text{decade}]$  with varying end year from 1998 to 2024, in which the red shaded envelope shows the bounds inside which a linear trend is not statistically significant according to a t-test at 95% confidence and (e,f) the 20 year-trends of sea ice area with varying end years but for each month of the year. The left column (a,c,e) shows the NSIDC sea index (Fetterer et al., 2017) and the right column (b,d,f) shows the OSISAF sea ice index (OSI-420, 2023).

across the CMIP5 and CMIP6 models, with the probability of a smaller-than-observed
2005-2024 trend varying from 0% and approximately 50%. Interestingly, the spread across
models for a given scenario is much larger than the spread across scenarios for a given
model. This is perhaps unsurprising because the scenarios diverge from each other later
than the 2020s (Notz & SIMIP Community, 2020), especially for the case of the CMIP6
forcing.

The central estimated value of approximately 20% doesn't change substantially if 218 models are selected following the criteria from Notz and SIMIP Community (2020), or 219 220 using models lying in the 66% range and 5-95% range estimates of the climate sensitivity (Sherwood et al., 2020), or using models according to their ability to simulate clima-221 tological sea ice conditions for the period 1979-1998 (Fig. 2a, column 2-4). Nor is the 222 central estimate substantially impacted if we assess the probability of a non-statistically 223 significant trend (Figure S3): this value is only slightly higher at approximately 25%. 224 Therefore from the multi-model perspective, what we have observed in the Arctic over 225 the past two decades is not a rare event, but rather one that should be expected to hap-226 pen with reasonable frequency. This result is insensitive to how models are sub-selected, 227 or to the metric of interest. This then raises the question of whether this pause in Arc-228 tic sea ice loss could continue and for how long. 229

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# **3.3** The observed pause in sea ice loss could foreseeably continue for another decade

To investigate how much longer this current pause is likely to last into the future, 232 we now examine those large ensemble members which do exhibit muted sea ice loss in 233 the period from 2005-2024 (Fig. 2a), and quantify how long the simulated slowdowns 234 persist in the future (Fig. 2b). In essence this produces the conditional probability es-235 timate of the 20-year pause extending further for each model and scenario. The multi-236 model average suggests that pauses in September sea ice loss for the period 2005-2024 237 have a 1 in 2 chance of persisting for a further five years, and a 1 in 3 chance of persist-238 ing for a further ten years (note, however, the considerable intermodel spread of  $\pm 25\%$ ). 239 On average, higher emissions scenarios tend to show slightly lower probabilities of sus-240 taining the muted pace of Arctic sea ice loss in the future, although the impact of this 241 is subtle and not consistent for every model. 242

On average, the sea ice area in ensemble members which simulate pauses in sea ice loss for the period 2005-2024 is 0.5 million km<sup>2</sup> larger in 2025 than in ensemble members from the same models and scenarios in which there is no pause (Fig. 2c). This source of predictability decays within a decade, and after that the September sea ice area in ensemble members with and without pauses are indistinguishable from each other.

It is important to highlight that to produce these estimates, we have limited the 248 model selection to only those models with at least five members which feature sea ice 249 pauses in 2004-2025, because to compute these probabilities in a meaningful way requires 250 that the remaining ensemble size is large enough. This may however, bias the results to-251 wards models with more ensemble members, which are more likely to include more mem-252 bers with slowdowns due to better sampling, and towards models which simulate slow-253 downs more frequently. The multi-model average of the probability of a sea ice loss pause 254 in this smaller subset of members is 28%, which is higher than the 20% estimated from 255 all models. If the observed slowdown is an inherently infrequent and rare event, then this 256 approach would overestimate the probability of it continuing, and underestimate how 257 anomalously high the Arctic sea ice cover is relative to the forced trend. However, of the 258 models which can reproduce the observed trends, over two-thirds of the available mod-259 els are included in this estimate. Our results are broadly consistent with the findings of 260 Swart et al. (2015), which showed that multi-decadal pauses longer than 20-years were 261 possible in the late 21st century under a medium emissions scenario. 262



Figure 2. (a) The percentage of ensemble members [%] for each ensemble that have 2005-2024 September sea ice area loss trends less than the observed value. The uncertainty estimate is calculated by monte carlo simulation with replacement. All simulations are shown on the left, with different selection criteria (that outlined in Notz and SIMIP Community (2020), the 5-95% and 66% range of ECS (Sherwood et al., 2020), and the climatological sea ice area applied on the right. (b) The conditional probability across each ensemble for the trends starting in 2005 to continue to be above -0.29 10<sup>6</sup> km<sup>2</sup>/decade for a given end year. (c) The ensemble-mean difference in September sea ice area [10<sup>6</sup> km<sup>2</sup>] between ensemble members with and without ice loss pauses over 2005-2024 for the period 2025-2050. For panels (b) and (c) only models and scenarios with at least five members with 2005-2024 trends above observed were included, with the black line shows the weighted average according to Equation 1. In all panels, each colour represents a model and each symbol represents a different forcing scenario.

Taken together, the wealth of available CMIP5 and CMIP6 simulations suggest it is possible, perhaps even likely, that the present slowdown in sea ice decline may continue for a further 5-10 years. If that were the case it may then imply the occurrence of an early ice-free Arctic is less likely than raw model output would suggest (Jahn et al., 2016; Arthun et al., 2020; Wang et al., 2021; England & Polvani, 2023; Jahn et al., 2024).

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#### 3.4 Climate models suggest an important role for climate variability

Whether the present slowdown persists in the future or not, one final question re-269 mains to be answered: is the recent pause a response to anthropogenic forcings alone, 270 or is there an important role for internal climate variability? When attempting to iso-271 late the forced component of any observed trend from internal variability, it is impor-272 tant to keep in mind that - assuming the model simulations faithfully capture a plau-273 sible reality - the observations are expected to have the same broad features as individ-274 ual ensemble members, i.e. that they are a combination of a forced trend plus one par-275 ticular realization of internal variability (although one doesn't expect the observations 276 to precisely match any one member). Given the well-established importance of internal 277 climate variability in Arctic sea ice trends (Kay et al., 2011; Swart et al., 2015; England 278 et al., 2019; Dörr et al., 2023), we next assess whether a change in the forced response 279 could also be substantially contributing to the observed slowdown in ice loss. 280

First, we show the forced September sea ice loss for the period 2005-2024 as esti-281 mated by the linear trend of the ensemble mean for each model and scenario (Fig. 3, hor-282 izontal axis). We find there are only two model/scenario combinations (GFDL-ESM2M 283 RCP8.5 and CESM2 SSP3-7.0) for which the forced trend is estimated to entirely ex-284 plain observed trends in ice loss (shown in Fig. 3, as grey vertical lines) with minimal 285 role for internal variability. Over 85% of the models we analyse here have a larger forced 286 sea ice loss for this period than observed (as they lie to the left of the observed trends), 287 implying that internal variability has acted to reduce the pace of ice loss. 288

Second, we ask: is there evidence that the forced response itself is slowing down 289 relative to the previous two decades? On the vertical axis of Figure 3, therefore, we plot 290 the ratio of the 2005-2024 forced trend to that of the preceding twenty years, 1986-2005, 291 for each model and scenario combination: a ratio of 1.0 indicates no change in the pace 292 of ice loss, > 1 indicates an acceleration and < 1 indicates a deceleration. Again we find 293 that only GFDL-ESM2M RCP8.5 and CESM2 SSP3-7.0 suggest that the reduction in 294 the forced trends accounts entirely for the observed slowdown. While the results from 295 all the other models agree that this observed pause is not entirely a forced response, the 296 remaining models are relatively evenly split (Fig. 3): roughly half the models suggest that 297 the forced sea ice loss trend has modestly decelerated over the past two decades relative 298 to the prior two decades, and roughly half suggest it has modestly accelerated (this is 200 especially clear if we disregard the models which are unable to capture the observed re-300 cent trends). 301

In summary then: while the modelling evidence is uncertain as to whether anthropogenic forcings - even in part - for the recent slowdown in Arctic sea ice loss, it is very likely that internal climate variability is contributing to the slowdown in an important way.

#### **306 4** Conclusion and Discussion

It is perhaps surprising that while global temperatures have risen rapidly, reaching record levels in the last few years, Arctic sea ice cover has shown no statistically significant decline over the past two decades. Nonetheless, analyzing two observational datasets and thousands of simulations from the CMIP5 and CMIP6 archives, we have established the following facts, which address the four questions raised in the introduction:



Figure 3. The ensemble mean trend in September sea ice area for the period 2005-2024 for each model and scenario (horizontal axis) versus the ratio of ensemble mean trends in September sea ice area for the periods 2005-2024 : 1986-2005 (vertical axis) with the black dotted line indicating a ratio of 1.0. Observational estimates of the 2005-2024 trend and the ratio of 2005-2024 : 1986-2005 trends are shown as grey lines (dashed line for NSIDC, solid line for OSISAF). Note that this does not imply that the observed trends are estimates of the forced response in the real climate system.

- 1. The pervasive slowdown of Arctic sea ice loss is robust across the choice of definitions, observational dataset, and season.
- 2. This observed pause in ice loss is simulated relatively frequently (with a 20% chance) 314 in climate models, and is thus to be expected even under high emission scenar-315 ios. 316
- 3. If model simulations are accurate, the recent pause may plausibly continue for an 317 additional five to ten years 318
  - 4. Nearly all models analysed suggest an important role for internal climate variability in slowing the anthropogenically-forced sea ice loss.

We now return to the question of the contribution of human influence versus in-321 ternal climate variability. If the slowdown is in fact a predominantly anthropogenically 322 forced episode, our results suggest that there must be either some shared missing forc-323 ing or common model deficiency in response to the standard forcing among the major-324 ity of the models. While the latter part is difficult to assess, one culprit for a missing forc-325 ing could be the increase in boreal forest fires, not incorporated in standard scenarios. 326 The recent study of Blanchard-Wrigglesworth et al. (2025) shows that incorporating re-327 cent biomass burning emissions into the CESM2 model leads to a rapid recovery of Septem-328 ber sea ice cover during the period of interest due to increased reflection in the North-329 ern Hemisphere of incoming shortwave radiation arising from the cloud response and aerosol 330 cloud interactions. However, due to specifics of the simulation of polar clouds in CESM2 331 (DeRepentigny et al., 2022; Zhu et al., 2022; Davis & Medeiros, 2024; England & Feldl, 332 2024), and a seemingly opposite result from similar experiments with a different climate 333 model (Zhong et al., 2024), further experiments with a wider range of models are needed 334 to understand the role changes in biomass burning have had on observed Arctic sea ice 335 trends. 336

Going forward, how can we use what we have learned about the recent pause in 337 Arctic sea ice loss? Firstly, if internal variability has played an important role then this 338 could provide a source of future predictability of Arctic climate change in the same man-339 ner as Yeager et al. (2015). And second, this period could be used as an out of sample 340 test in future climate model evaluation – similar to the early and middle periods of the 341 20th century (Flynn et al., 2023; Bianco et al., 2024; Chen & Dai, 2024). However, over-342 all this study is a reminder that we should be humble about multi-decadal predictions 343 of the climate system, especially in highly variable regions such as the Arctic. Standing in 2007 or 2012 after having experienced another year of record loss and listening to as-345 sessments that climate models are flawed in their ability to reproduce the rapid loss of 346 Arctic sea ice (Stroeve et al., 2007), it would take a rather brave person to have predicted 347 that a sustained slowdown in ice loss was around the corner, although, as we have shown, 348 and many have found before (Kay et al., 2011; Swart et al., 2015; R. Zhang, 2015), this 349 is entirely consistent with what climate models simulate. 350

#### **Open Research Section** 351

All CMIP5 and CMIP6 data analysed in this study is publicly available to down-352 load from the Earth System Federation Grid at https://aims2.llnl.gov/search. 353

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#### 360 References

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- 361Arthun, M., Onarheim, I., Dörr, J., & Eldevik, T.(2020).The seasonal and362regional transition to an ice-free Arctic.Geophysical Research Letters,36348 (e2020GL090825).doi: 10.1029/2020GL0909825
- Baxter, I., Ding, Q., Schweiger, A., L'Heureux, M., Baxter, S., Wang, T., ... Lu, J.
   (2019). How tropical Pacific surface cooling contributed to accelerated sea ice melt from 2007 to 2012 as ice is thinned by anthropogenic forcing. *Journal of Climate*, 32, 8583-8602. doi: 10.1175/JCLI-D-18-0783.1
- Bianco, E., Blanchard-Wrigglesworth, E., Materia, S., Ruggieri, P., Iovino, D.,
  - & Masina, S. (2024). CMIP6 models underestimate Arctic sea ice loss during the early twentieth-century warming, despite simulating large low-
  - frequency sea ice variability. Journal of Climate, 37, 6305-6321. doi: 10.1175/JCLI-D-23-0647.1
- Blanchard-Wrigglesworth, E., DeRepentigny, P., & Frierson, D. (2025). Increasing
   boreal fires reduce future global warming and sea ice loss. *Preprint*. doi: 10
   .21203/rs.3.rs-5486231/v2
- Chen, X., & Dai, A. (2024). Quantifying contributions of external forcing and internal variability to Arctic warming during 1900-2021. *Earth's Future*, *12*(e2023EF003734). doi: 10.1029/2023EF003734
- Copernicus, E. (2024, April). *Climate indicators.* Retrieved 16/02/2025, from https://climate.copernicus.eu/climate-indicators
  - Davis, I., & Medeiros, B. (2024). Assessing CESM2 clouds and their response to climate change using cloud regimes. Journal of Climate, 37, 2965-2985. doi: 10 .1175/JCLI-D-23-0337.1
- DeRepentigny, P., Jahn, A., Holland, M., Kay, J., Fasullo, J., Lamarque, J., ... Barrett, A. (2022). Enhanced simulated early 21st century Arctic sea ice loss due to CMIP6 biomass burning emissions. *Science Advances*, 8(30). doi: 10.1126/sciadv.abo2405
  - Deser, C., & Phillips, A. (2021). Defining the internal component of Atlantic Multidecadal Variability in a changing climate. Geophysical Research Letters, 48 (e2021GL095023). doi: 10.1029/2021GL095023
  - Ding, Q., Schweiger, A., L'Heureux, M., Steig, E., Battisti, D., Johnson, N., ...
  - Baxter, I. (2018). Fingerprints of internal drivers of Arctic sea ice loss in observations and model simulations. *Nature Geoscience*, 12, 28-33. doi: 10.1038/s41561-018-0256-8
- Dörr, J., Bonan, D., Arthun, M., Svendsen, L., & Wills, R. (2023). Forced and inter nal components of observed Arctic sea-ice changes. *The Cryosphere*, 17, 4133 4153. doi: 10.5194/tc-17-4133-2023
- England, M. (2021). Are multi-decadal fluctuations in arctic and antarctic surface
   temperatures a forced response to anthropogenic emissions or part of internal
   climate variability? *Geophysical Research Letters*, 48 (e2020GL090631). doi:
   10.1029/2020GL090631
- England, M., Eisenman, I., Lutsko, N., & Wagner, T. (2021). The recent emergence
   of Arctic Amplification. *Geophysical Research Letters*, 48 (e2021GL094086).
   doi: 10.1029/2021GL094086
- England, M., & Feldl, N. (2024). Robust polar amplification in ice-free climates
   relies on ocean heat transport and cloud radiative effects. Journal of Climate,
   37, 2179-2197. doi: 10.1175/JCLI-D-23-0151.1
- England, M., Jahn, A., & Polvani, L. (2019). Nonuniform contribution of internal
   variability to recent Arctic sea ice loss. Journal of Climate, 32, 4039-4053. doi:
   10.1175/JCLI-D-18-0864.1
- England, M., & Polvani, L. (2023). The Montreal Protocol is delaying the occurrence of the first ice-free Arctic summer. *PNAS*, 120 (e2211432120). doi: 10
  .1073/pnas.2211432120

Fetterer, F., Knowles, K., Meier, W., Savoie, M., & Windnagel, A. (2017). Sea ice 414 index. (q02135, version 3). [dataset]. National Snow and Ice Data Center. doi: 415 10.7265/N5K072F8 416 Flynn, C., Huusko, L., Modak, A., & Mauritsen, T. (2023).Strong aerosol cool-417 ing alone does not explain cold-biased mid-century temperatures in CMIP6 418 models. Atmospheric Chemistry and Physics, 23, 15,121-15,133. doi: 419 10.5194/acp-23-15121-2023 420 Francis, J., & Wu, B. (2020). Why has no new record-minimum Arctic sea-ice extent 421 occurred since september 2012? Environmental Research Letters, 15(114034). 422 doi: 10.1088/1748-9326/abc047 423 Fyfe, J., von Salzen, K., Gillett, N., Arora, V., Flato, G., & McConnell, J. (2013).424 One hundred years of Arctic surface temperature variation due to anthro-425 pogenic influence. Scientific Reports, 3(2645). doi: 10.1038/srep02645 426 Gagne, M., Fyfe, J., Gillett, N., Polyakov, I., & Flato, G. (2017).Aerosol-driven 427 increase in Arctic sea ice over the middle of the twentieth century. Geophysical 428 Research Letters, 44, 7338-7346. doi: 10.1002/2016GL071941 429 Hansen, J., Kharecha, P., Sato, M., Tselioudis, G., Kelly, J., Bauer, S., ... Pokela, 430 Α. (2025).Global warming has accelerated: Are the United Nations and 431 the public well-informed? Environment: Science and Policy for Sustainable 432 Development, 67, 6-44. doi: 10.1080/00139157.2025.2434494 433 IPCC. (2021).Climate change 2021: The physical science basis. Contribution of 434 working group I to the sixth assessment report of the Intergovernmental Panel 435 on Climate Change. In V. Masson-Delmotte et al. (Eds.), (p. 3-32). Cambridge 436 University Press. 437 (2024).Jahn, A., Holland, M., & Kay, J. Projections of an ice-free Arctic Ocean. 438 Nature Reviews Earth and Environment, 5, 167-176. doi: 10.1038/s43017-023 439 -00515-9440 Jahn, A., Kay, J., Holland, M., & Hall, D. (2016). How predictable is the timing of a 441 summer ice-free Arctic? Geophysical Research Letters, 43, 9113-9120. doi: 10 442 .1002/2016GL070067 443 Kay, J., Holland, M., & Jahn, A. (2011). Inter-annual to multi-decadal Arctic sea ice 444 extent trends in a warming world. Geophysical Research Letters, 38(L15708). 445 doi: 10.1029/2011GL048008 446 Kerr, R. (2000). A North Atlantic climate pacemaker for the centuries. Science. 447 288, 1984-1985. doi: 10.1126/science.288.5473.1984 448 Kwok, R. (2018). Arctic sea ice thickness, volume, and multiyear ice coverage: losses 449 and coupled variability (1958–2018). Environmental Research Letters, 13, 450 105005. doi: 10.1088/1748-9326/aae3ec 451 Lee, Y., & Liu, W. (2023).The weakened Atlantic Meridional Overturning Cir-452 culation diminishes recent Arctic sea ice loss. Geophysical Research Letters, 453 50(e2023GL105929). doi: 10.1029/2023GL105929 454 Anthropogenic aerosol forcing under Lund, M., Myhre, G., & Samset, B. (2019).455 the Shared Socioeconomic Pathways. Atmospheric Chemistry and Physics, 19, 456 13827-13839. doi: 10.5194/acp-19-13827-2019 457 Manshausen, P., Watson-Parris, D., Christensen, M., Jalkanen, J., & Stier, P. 458 (2022).Invisible ship tracks show large cloud sensitivity to aerosol. Nature. 459 610, 101–106. doi: 10.1038/s41586-022-05122-0 460 Mantua, N., & Hare, S. (2002). The Pacific Decadal Oscillation. Journal of Oceanog-461 raphy, 58, 35-44. doi: 10.1023/A:1015820616384 462 Maslowski, W., Kinney, J., Higgins, M., & Roberts, A. (2012). The future of Arctic 463 sea ice. Annual Review of Earth and Planetary Sciences, 40, 625-654. doi: 10 464 .1146/annurev-earth-042711-105345 465 Meehl, G., Chung, C., Arblaster, J., Holland, M., & Bitz, C. (2018).Tropical 466 decadal variability and the rate of Arctic sea ice decrease. Geophysical Re-467 search Letters, 45, 11,326-11,333. doi: 10.1029/2018GL079989 468

469	Meehl, G., Senior, C., Eyring, V., Lamarque, J., Stouffer, R., Taylor, K., & Schlund,
470	M. (2020). Context for interpreting equilibrium climate sensitivity and tran-
471	sient climate response from the CMIP6 Earth system models. Science Ad-
472	vances, 6 (eaba1981). doi: 10.1126/sciadv.aba1981
473	Meier, W., & Stewart, J. (2019). Assessing uncertainties in sea ice extent climate
474	indicators. Environmental Research Letters, 14(035005). doi: 10.1088/1748
475	-9326/aaf52c
476	Merchant, C., Allan, R., & Embury, O. (2025). Quantifying the acceleration of mul-
477	tidecadal global sea surface warming driven by Earth's energy imbalance. En-
478	vironmental Research Letters, 20(024037). doi: 10.1088/1748-9326/adaa8a
479	Nafaji, M., Zwiers, F., & Gillett, N. (2015). Attribution of Arctic temperature
480	change to greenhouse-gas and aerosol influences. Nature Climate Change, 5.
481	246-249. doi: 10.1038/nclimate2524
482	Notz, D., & SIMIP Community. (2020). Arctic sea ice in CMIP6. Geophysical Re-
483	search Letters, 47(e2019GL086749). doi: 10.1029/2019GL086749
484	Notz, D., & Stroeve, J. (2016). Observed Arctic sea-ice loss directly follows anthro-
485	pogenic CO2 emission. <i>Science</i> , 354, 747-750, doi: 10.1126/science.aag2345
196	Onarheim I Arthun M Teigen S Eik K & Steele M (2024) Recent
400	thickening of the Barents Sea ice cover Geonbusical Research Letters
401	51 (e2024GL108225) doi: 10 1029/2024GL108225
400	OSI-420 (2023) Osi saf sea ice inder 1978-onwards version 2.2 [dataset] EU-
409	METSAT Ocean and Sea Ice Satellite Application Facility, doi: 10.15770/EUM
490	SAF OSI 0022
491	Parkinson C & Comiso I (2013) On the 2012 record low Arctic sea ice cover:
492	Combined impact of preconditioning and an August storm <i>Ceonbusical Re-</i>
493	search Letters 10 1356-1361 doi: 10 1002/grl 50349
494	Bantanan M. Karpachko A. Lipponan A. Huyarinan O. Buostaanaja K.
495	Vibra T & Laaksonen $\Lambda$ (2022) The Arctic has warmed nearly four times
496	faster than the globe since 1070 Communications Earth and Environment
497	3(168) doi: 10.1038/s/3207-022-00/08-3
498	Samset B. Zhou, C. Fuglestvedt I. Lund M. Marotzka, I. & Zelinka, M. (2023)
499	Steady global surface warming from 1073 to 2022 but increased warming
500	rate after 1990 Communications Earth and Environment /(400) doi:
501	$10 \ 1038 \ s 43247.023.01061.4$
502	Schweiger A Lindsey R Zhang I Steele M Stern H & Kwok R (2011) Un-
503	certainty in modeled Arctic sea ice volume Journal of Geophysical Research
504	116(C00D06) doi: 10.1029/2011.IC007084
505	Shen Z. Duan A. Zhou W. Peng V. & Li I. (2024). Reconciling roles of ex-
500	ternal forcing and internal variability in Arctic sea ice change on different time
507	scales Lowrnal of Climate 37 3577-3591 doi: 10.1175/JCLI-D-23-0280.1
500	Sherwood S Webb M Annan I Armour K Forster P Hargreaves I
509	Zelinka M (2020) An assessment of Earth's climate sensitivity using mul-
510	tiple lines of evidence <i>Review of Geophysics</i> 58(e2019BG000678) doi:
511	10 1029/2019BG000678
512	Siew P. Wu V. Ting M. Zheng C. Ding O. & Seager B. (2024) Signif-
515	icant contribution of internal variability to recent Barents-Kara sea ice
514	loss in winter <u>Communications Earth and Environment</u> 5(411) doi:
515	10 1038/s43247_024_01582_6
510	Stroeve I Holland M Mejer W Scambos T & Sarrozo M (2007) Arctic
517	sea ice decline: faster than forecast. <i>Geophysical Research Letters</i> 9/(L00501)
510	doi: 10.1029/2007GL029703
213	Stroeve I & Notz D (2018) Changing state of Arctic con ico across all concord
520	Environmental Research Letters 12(103001) doi: 10.1082/1748.0296/204056
521	Stroeve I Serreze M Drobot S Coerboad S Holland M Madanika I
522	Scombos T (2011) Arctia con ico extent plummete in 2007 Eco 20 12 14
523	Scambos, 1. (2011). Arctic sea ice extent plummets in 2007. $Eos, \delta 9, 13-14$ .

524	doi: 10.1029/2008EO020001
525	Swart, N., Fyfe, J., Hawkins, E., Kay, J., & Jahn, A. (2015). Influence of internal
526	variability on arctic sea-ice trends. Nature Climate Change, 5, 86-89. doi: 10
527	.1038/nclimate2483
528	Szopa, S., Balkanski, Y., Schulz, M., Bekki, S., Cugnet, D., Fortems-Cheiney, A.,
529	Dufresne, J. (2013). Aerosol and ozone changes as forcing for climate
530	evolution between 1850 and 2100. Climate Dynamics, 40, 2223-2250. doi:
531	10.1007/s00382-012-1408-y
532	Wadhams, P. (2016). A farewell to ice. Penguin Publishing.
533	Walsh, J., Fetterer, F., Scott, J., & Chapman, W. (2017). A database for depicting
534	Arctic sea ice variations back to 1850. Geographical Review, 107, 89-107. doi:
535	10.1111/j.1931-0846.2016.12195.x
536	Wang, B., Zhou, X., Ding, Q., & Liu, J. (2021). Increasing confidence in project-
537	ing the Arctic ice-free year with emergent constraints. Environmental Research
538	Letters, $16(094016)$ . doi: $10.1088/1748-9326/ac0b17$
539	Yeager, S., Karspeck, A., & Danabasoglu, G. (2015). Predicted slowdown in the rate
540	of Atlantic sea ice loss. Geophysical Research Letters, 42, 10,704-10,713. doi:
541	$10.1002/2015 { m GL}065364$
542	Yoshioka, M., Grosvenor, D., Booth, B., Morice, C., & Carslaw, K. (2024). Warming
543	effects of reduced sulfur emissions from shipping. Atmospheric Chemistry and
544	<i>Physics</i> , $24$ , 13681–13692. doi: 10.5194/acp-24-13681-2024
545	Zhang, J. (2021). Recent slowdown in the decline of Arctic sea ice volume under
546	increasingly warm atmospheric and oceanic conditions. Geophysical Research
547	<i>Letters</i> , 48(e2021GL094780). doi: 10.1029/2021GL094780
548	Zhang, J., Lindsay, R., Schweiger, A., & Steele, M. (2013). The impact of an intense
549	summer cyclone on 2012 Arctic sea ice retreat. Geophysical Research Letters,
550	40,720-726. doi: 10.1002/gr1.50190
551	Zhang, R. (2015). Mechanisms for low-frequency variability of summer Arctic sea ice
552	extent. PNAS, 112, 4570-4575. doi: 10.1073/pnas.1422296112
553	Zhong, Q., Shutgens, N., Veraverbeke, S., & van der Werf, G. (2024). Increasing
554	aerosol emissions from boreal biomass burning exacerbate Arctic warming. Na-
555	ture Change, 14, 1275-1281. doi: 10.1038/s41558-024-02176-y
556	Linu, J., Otto-Biesner, B., Brady, E., Gettelman, A., Bacmeister, J., Neale, R.,
557	ray, J. (2022). LGW paleoclimate constraints inform cloud parameteriza-
558	tions and equilibrium chinate sensitivity in CESM2. Journal of Advances in Modeling Forth Systems, 17(20021MS002776), doi: 10.1020/2021MS002776
559	<i>Modeling Earth Systems</i> , 14 (e2021MS002776). doi: 10.1029/2021MS002776

# Supporting Information for "Surprising, but not unexpected, multi-decadal pause in Arctic sea ice loss"

M. R. England<sup>1</sup>, L. M. Polvani<sup>2,3</sup>, J. Screen<sup>1</sup>, and A. C. Chan<sup>1</sup>

 $^{1}\mathrm{Department}$  of Mathematics and Statistics, University of Exeter, UK

 $^{2}\mathrm{Lamont}\text{-}\mathrm{Doherty}$ Earth Observatory, Columbia University, New York, USA

<sup>3</sup>Department of Applied Physics and Applied Mathematics, Columbia University, New York, USA

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## References

- Boucher, O., Servonnat, J., & co-authors. (2020). Presentation and evaluation of the IPSL-CM6A-LR climate model. Journal of Advances in Modeling Earth Systems, 12(e2019MS002010). doi: 10.1029/2019MS002010
- Burger, F., Terhaar, J., & Frölicher, T. (2022). Compound marine heatwaves and ocean acidity extremes. Nature Communications, 13(4722). doi: 10.1038/s41467-022-32120 -7

Fasullo, J., Golaz, J., Caron, J., Rosenbloom, N., Meehl, G., Strand, W., ... Bartoletti,
T. (2024). An overview of the E3SM version 2 Large Ensemble and a comparison to other E3SM and CESM Large Ensembles. *Earth System Dynamics*, 15, 367-386. doi: 10.5194/esd-15-367-2024

:

- Kay, J., Deser, C., Phillips, A., Mai, A., Hannay, C., Strand, G., ... Vertenstein, M. (2015). The Community Earth System Model (CESM Large Ensemble project: A community resource for studying climate change in the presence of internal climate variability. *Bulletin of the American Meteorological Society*, 96, 1333-1349. doi: 10.1175/BAMS-D-13-00255.1
- Kirchmeier-Young, M., Zwiers, F., & Gillett, N. (2017). Attribution of extreme events in Arctic sea ice extent. Journal of Climate, 30, 553-571. doi: 10.1175/JCLI-D-16 -0412.1
- Maher, N., Milinski, S., Suarez-Gutierrez, L., Botzet, M., Kornblueh, L., Kroger, J., ... Marotzke, J. (2019). The Max Planck Institute Grand Ensemble: Enabling the exploration of climate system variability. *Journal of Advances in Modeling Earth Systems*, 11, 2050-2069. doi: 10.1029/2019MS001639
- Olonscheck, D., Suarez-Gutierrez, L., Milinski, S., Beobide-Arsuaga, G., Baerh, J., Frob,
  F., ... Brune, S. (2023). The new Max Planck Institute Grand Ensemble with
  CMIP6 forcing and high-frequency model output. Journal of Advances in Modeling
  Earth Systems, 15(e2023MS003790). doi: 10.1029/2023MS003790
- Rodgers, K., Lee, S., Rosenbloom, N., Timmermann, A., Danabasoglu, G., Deser, C., ... Yeager, S. (2021). Ubiquity of human-induced changes in climate variability. *Earth*

System Dynamics, 12, 1393-1411. doi: 10.5194/esd-12-1393-2021

- Sellar, A., Jones, C., & co-authors. (2019). UKESM1: Description and evaluation of the UK Earth System Model. Journal of Advances in Modeling Earth Systems, 11, 4513-4558. doi: 10.1029/2019MS001739
- Sun, L., Alexander, M., & Deser, C. (2018). Evolution of the global coupled climate response to Arctic sea ice loss during 1990-2090 and its contribution to climate change. *Journal of Climate*, 31, 7823-7843. doi: 10.1175/JCLI-D-18-0134.1
- Swart, N., Cole, J., Kharin, V., Lazare, M., Scinocca, J., Gillett, N., ... Winter, B. (2019). The Canadian Earth System Model version 5 (CanESM5.0.3). Geoscientific Model Development, 12, 4823-4873. doi: 10.5194/gmd-12-4823-2019
- Tatebe, H., Ogura, T., Nitta, T., Komuro, Y., Ogochi, K., Takemura, T., ... Kimoto, M. (2019). Description and basic evaluation of simulated mean state, internal variability, and climate sensitivity of MIROC6. *Geoscientific Model Development*, 12, 2727-2765. doi: 10.5194/gmd-12-2727-2019
- Wyser, K., Koenigz, T., Fladrich, U., Fuentes-Franco, R., Karami, M., & Kruschke, T. (2021). The SMHI Large Ensemble (SMHI-LENS) with EC-Earth3.3.1. Geoscientific Model Development, 14, 4781-4796. doi: 10.5194/gmd-14-4781-2021
- Ziehn, T., Chamberlain, M., Law, R., Lenton, A., Bodman, R., Dix, M., ... Srbinovsky,
  J. (2020). The Australian Earth System Model: ACCESM-ESM1.5. Journal of Southern Hemisphere Earth Systems Science, 80, 193-214. doi: 10.1071.ES19035



**Figure S1.** Timeseries of September sea ice area (top row) and sea ice extent (bottom row) for the NSIDC (left column) and OSISAF (right column) sea ice indices. The 95% confidence interval for the 20-year linear trend 2005-2024 is shown in the shading.





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Figure S2. Timeseries of PIOMAS-simulated Arctic sea ice volume anomaly for the period 1979-2024. Anomalies are calculated as the departure from the long term 1979-2024 average, with daily anomalies shown in the grey and annual anomalies shown in the black. The 95% confidence interval for the 15-year linear trend 2010-2024 is shown in the shading.



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Figure S3. Same as in Figure 2 but for the percentage of members for which 2005-2024 September Arctic sea ice area trends are not statistically significant at 95% confidence.

Table S1. Detai	ls of the large	e ensemble simulations analysed in this study.	
Model	Generation	Scenarios (members)	Reference
CESM1	CMIP5	RCP4.5 (15), RCP8.5 (40)	(Kay et al., 2015)
GFDL-CM3	CMIP5	RCP8.5 (20)	(Sun et al., 2018)
GFDL-ESM2M	CMIP5	RCP8.5 (30)	(Burger et al., 2022)
CanESM2	CMIP5	RCP8.5(50)	(Kirchmeier-Young et al., 2017)
MPI-ESM-LR	CMIP5	RCP2.6, 4.5, and 8.5 (100)	(Maher et al., 2019)
ACCESS-ESM1-5	CMIP6	SSP1-2.6, 2-4.5, 3-7.0, and 5-8.5 (40)	(Ziehn et al., 2020)
CESM2	CMIP6	SSP2-4.5 (16), SSP3-7.0 (50 <sup>a</sup> )	(Rodgers et al., 2021)
CanESM5	CMIP6	SSP2-4.5, 3-7.0, and 5-8.5 (10)	(Swart et al., 2019)
EC-Earth 3	CMIP6	SSP1-1.9, 1-2.6, 2-4.5, 3-7.0, 4-3.4, 4-6.0, and 5-8.5 (50)	(Wyser et al., 2021)
IPSL CM6A	CMIP6	SSP2-4.5 and 3-7.0 (11)	(Boucher et al., 2020)
UKESM1-0-LL	CMIP6	SSP1-2.6 (16), SSP2-4.5 (15), SSP3-7.0 (16)	(Sellar et al., $2019$ )
MIROC6	CMIP6	SSP1-1.9, SSP1-2.6, 2-4.5, 3-7.0, and 5-8.5 (50)	(Tatebe et al., 2019)
MPI-ESM1.2	CMIP6	SSP1-1.9, 1-2.6, 2-4.5, 3-7.0, and 5-8.5 (50)	(Olonscheck et al., 2023)
E3SM-2-0	CMIP6	SSP3-7.0 (21)	(Fasullo et al., 2024)
<sup>a</sup> This is the mer	mbers with o	r without biomass.	